

1 **Topography-based estimation of evapotranspiration at** 2 **high altitudes in semi-arid regions** 3 Badreddine Sebbar^{1,2}, Olivier Merlin², Saïd Khabba^{1,3}, Vincent Simonneaux², Marine 4 Bouchet², and Abdelghani Chehbouni^{1,2, 4} $\frac{5}{6}$ ¹ Center for Remote Sensing Applications, Mohammed VI Polytechnic University, Benguerir, Morocco $\begin{array}{c} 7 \\ 8 \end{array}$ ²Centre d′Etudes Spatiales de la Biosphère (CESBIO), Université de Toulouse, CNES, CNRS, 8 IRD, UPS, Toulouse, France
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14 **Abstract.** Assessing the surface water balance of mountains is a real challenge given notably 15 the extreme variability of meteorological conditions and the sparsity of in situ monitoring. While
16 mountains are recognized as water towers feeding the surrounding plains, there is only unconsol-16 mountains are recognized as water towers feeding the surrounding plains, there is only unconsol-
17 idated knowledge about the individual water balance components especially the evanotranspira-17 idated knowledge about the individual water balance components especially the evapotranspira-
18 ion (ET). Satellites Land Surface Temperature (LST) along with air temperature (Ta) and global 18 tion (ET). Satellites Land Surface Temperature (LST) along with air temperature (Ta) and global 19 solar radiation (Rg) can be used to assess the energy budget and provide a reasonable estimation
20 of instantaneous ET. Nevertheless, over mountains the Ta and Rg respectively undergo strong 20 of instantaneous ET. Nevertheless, over mountains the Ta and Rg respectively undergo strong
21 topographical changes due to elevation and sun exposure effects. Moreover, upscaling the instan-21 topographical changes due to elevation and sun exposure effects. Moreover, upscaling the instan-
22 taneous ET to its daily value is expected to be uncertain in mountains as the evaporative fraction 22 taneous ET to its daily value is expected to be uncertain in mountains as the evaporative fraction
23 (EF, defined as the ratio of ET to available energy ratio) of a given pixel can no longer be con-23 (EF, defined as the ratio of ET to available energy ratio) of a given pixel can no longer be con-24 sidered constant during daytime until proven otherwise. In this context, this contribution focuses 25 on a topography-based estimation of ET using the Two Source Energy Balance (TSEB) model. 26 We also examine the variability of hourly and daily EF estimates through both satellite and in
27 situ monitoring. An eddy covariance tower was installed at 3850 m.a.s.l over the High Atlas 27 situ monitoring. An eddy covariance tower was installed at 3850 m.a.s. l over the High Atlas
28 Mountains in central Morocco and has been operating since September 2020 to present. 30 m 28 Mountains in central Morocco and has been operating since September 2020 to present. 30 m 29 resolution LST is derived from thermal data collected by Landsat-7, 8 and 9 on clear sky days.
 29 Rg is estimated at the Landsat (30 m) resolution from the SRTM's Digital Elevation Model 30 Rg is estimated at the Landsat (30 m) resolution from the SRTM's Digital Elevation Model 31 (DEM) and two different topography-based approaches: a physically based model (DART) and 32 a simplified semi-empirical model. The 9 km resolution ERA5-Land's air temperature product is
33 spatialized at the same (30 m) resolution by applying the environmental lanse rate (ELR) re-33 spatialized at the same (30 m) resolution by applying the environmental lapse rate (ELR) retrieved at the Landsat overpass time over a 9 Km² area including the eddy covariance tower. trieved at the Landsat overpass time over a 9 Km² area including the eddy covariance tower. 35 Satellite-derived estimates of ET and EF are compared to instantaneous station measurements for 36 three and nine dates in 2020 and 2021 respectively. The variability during daytime of the in-situ
37 EF is also assessed to evaluate the potential for upscaling instantaneous remotely sensed ET to a 37 EF is also assessed to evaluate the potential for upscaling instantaneous remotely sensed ET to a daily scale.

39 **Keywords:** Evapotranspiration, Topography, Evaporative Fraction, Remote Sensing, Energy-balance.

41 **1 Introduction**

42 The estimation of evapotranspiration (ET) in mountainous areas using temperature-43 based models that rely on remotely sensed Land Surface Temperature (LST) is chal lenging due to the influence of topographical effects. The elevation-dependent temper-45 ature (Ta) and incoming radiation (Rg) from adjacent surfaces make it difficult to spa- tially distribute the necessary inputs. To overcome this challenge, a new methodologi- cal approach is proposed in this study. This approach considers these effects prior to estimating ET in mountainous regions.

2 Materials and Methods

 The study area is the Rheraya sub-basin located in the High Atlas Mountains in central Morocco. The area has a semi-arid climate, and its elevation varies from about 1000 to 4127 meters. The high-altitude regions are characterized by low temperatures, rugged terrain, and sparse vegetation cover.

 Malbeteau et al. [1] normalized LST for topographical effects. Dynamic environmental lapse rate (ELR) was physically inverted through an energy balance model, whereas Rg was spatialized based on the DART model [2]. Ta is then spatialized following the expression:

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T_{pixel} = T_{station} + ELR(E_{pixel} - E_{station})
$$
 (1)

 We adapted a similar methodological approach through replacing in-situ measurements with ERA5-Land meteorological data. Rg was estimated using an alternative approach based on direct radiation [3] instead of DART (which is potentially very precise but difficult to apply over large areas).

 ET is estimated as a residual of the two-source energy balance (TSEB) model. The latter calculates the energy balance of the soil-canopy-atmosphere continuum, where transpiration is initially determined by the Priestley-Taylor equation [4]. In our case the model is triggered by the aforementioned meteorological forcing and the LST.

3 Results

3.1 Ta and Rg spatialization

 Figure 1 depicts scatterplots of simulated Rg and Ta compared to in-situ observations using the Samani [3] and ELR methods, respectively. In addition, it compares measured and simulated Rn based on the resulting Rg and Ta, as well as LST. Rg, Ta, and Rn each have RMSE values of 49.80 W/m2, 1.9 °C, and 43.30 W/m2 correspondingly.

0.80, 0.85, and 0.76 are the R2 values for the same variables in the same sequence.

 Fig. 1. Scatterplots of measured vs modeled Rg, Ta and Rn at the top-left corner, top-right corner, and the bottom respectively.

3.2 EF daily variation

 Figure 2 compares the model-based instantaneous EF estimate at the satellite over- pass to hourly and daily measurements. The satellite's midday overpass gives an inter- esting configuration; the projected instantaneous EF values are fairly close to the daily observed EF. The highest recorded bias is approximately 5%, which corresponds to the 7th of August 2021, when the observed EF around sunset rises abruptly due to instru-mentation faults.

 Fig. 2. Instantaneous remotely sensed (red dot) vs daily/hourly (green line/blue line) in-situ measured EF.

3.3 ET estimation

 Taking topographic effects into account in Figure 3 results in an RMSE of 27.08 W/m2 87 and an R2 of 0.81 for the simulated ET; for Rn, these values are 42.73 W/m2 and 0.67; 88 and for the sensible heat flux (H), they are 56.43 W/m2 and 0.79. When topography is considered, the findings mostly agree with the observations, except for the conductive 90 flux (G), which was approximated as a percentage of Rn (RMSE = 73.26 W/m2, R2 = 0.12).

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 Fig. 3. Comparison of topographically-affected and non-topographically-affected modeled energy balance fluxes with Eddy-covariance observations.

4 Discussion

 The results of our study indicate that the Rg, Ta, and Rn spatialization techniques used in the TSEB-PT model are highly effective in predicting instantaneous fluxes, with val- ues that closely observations. Our analysis also revealed that the instantaneous EF could potentially be used to extrapolate instantaneous to daily ET values, particularly when the satellite overpass occurs around midday. This finding is consistent with previous studies that have demonstrated the utility of the EF in estimating daily energy fluxes over flat regions (e.g., Crago, R. D., 1996; Hoedjes et al., 2008). Furthermore, our study highlights the significant influence of topography on modeled fluxes, as evidenced by strong correlations between modeled and measured fluxes after daily extrapolation. These results are consistent with previous studies that have shown the importance of incorporating topographic information in energy flux estimation models (e.g., Rana et al., 2007; Hao et al., 2021). Overall, our study contributes to the growing body of liter- ature on energy flux estimation in terrestrial ecosystems and supports the use of the TSEB-PT model in accurately predicting energy fluxes in various meteorological and topographic conditions.

5 Conclusions

 The estimation of ET over mountainous terrain is the primary emphasis of this work. The removal of topographic influences is essential for achieving an accurate estimate of the spatial variation in ET values. When topographic effects are considered, the TSEB-PT model exhibits satisfactory performance. As a perspective, our aim is to es- timate ET by applying a more straightforward contextual model, such as the LST-VI model while considering topographic normalization of LST.

References

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